



# *PICES Sixteenth Annual Meeting*

## **S2**

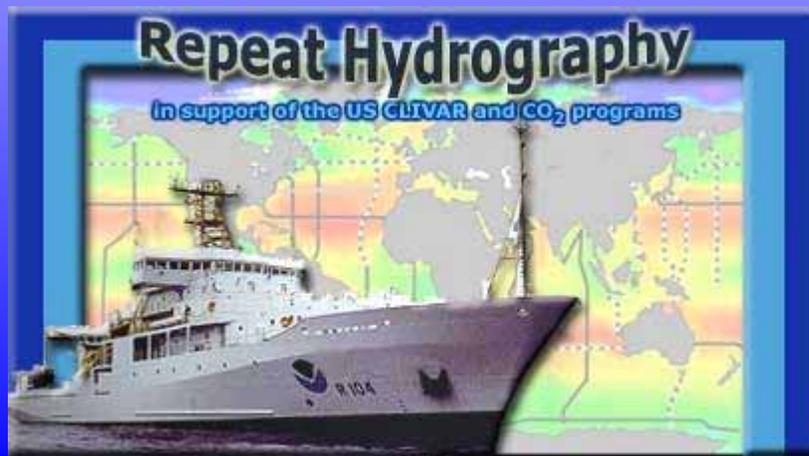
**BIO/POC Topic Session**

**Decadal changes in carbon biogeochemistry in the North Pacific**

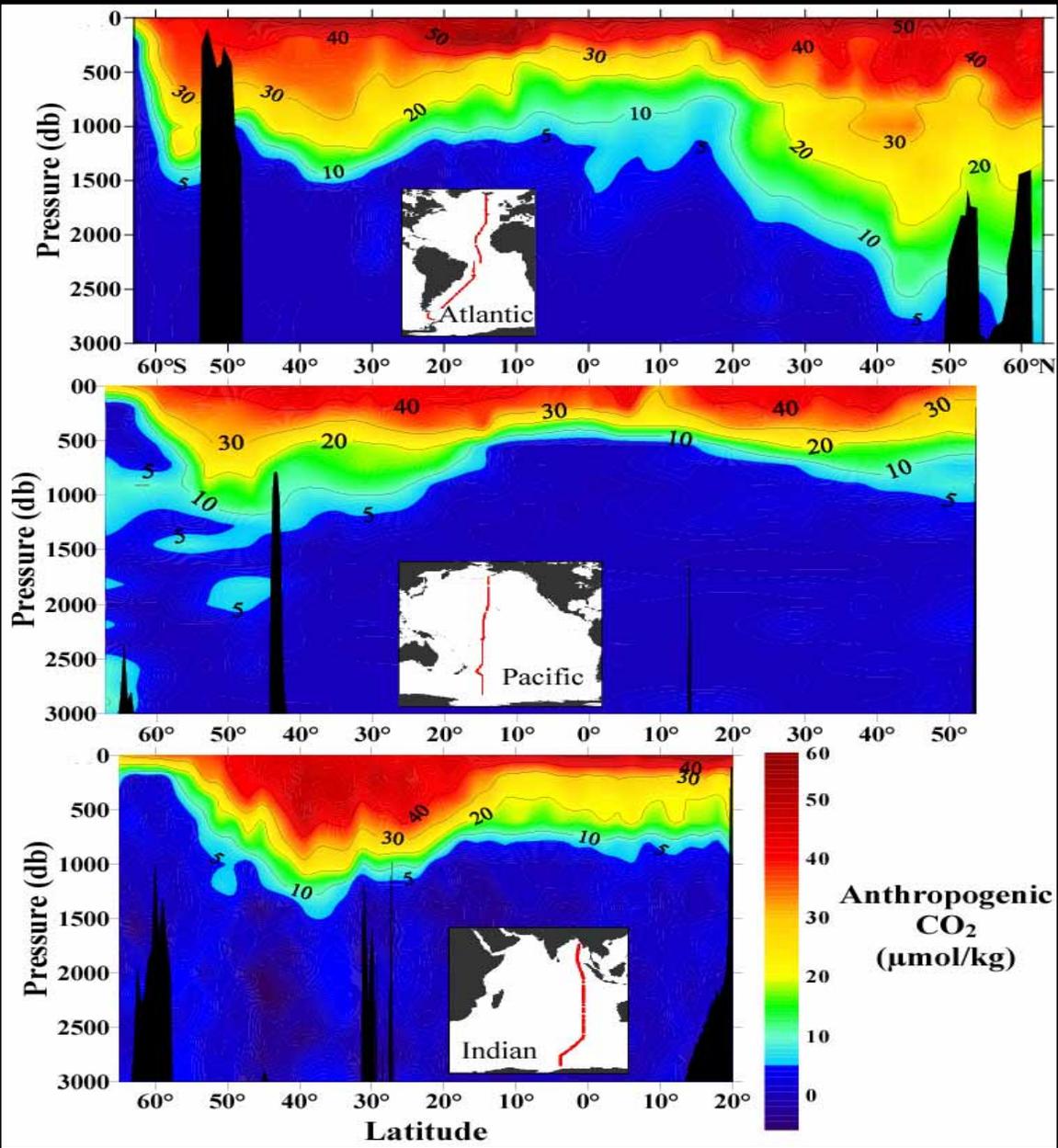
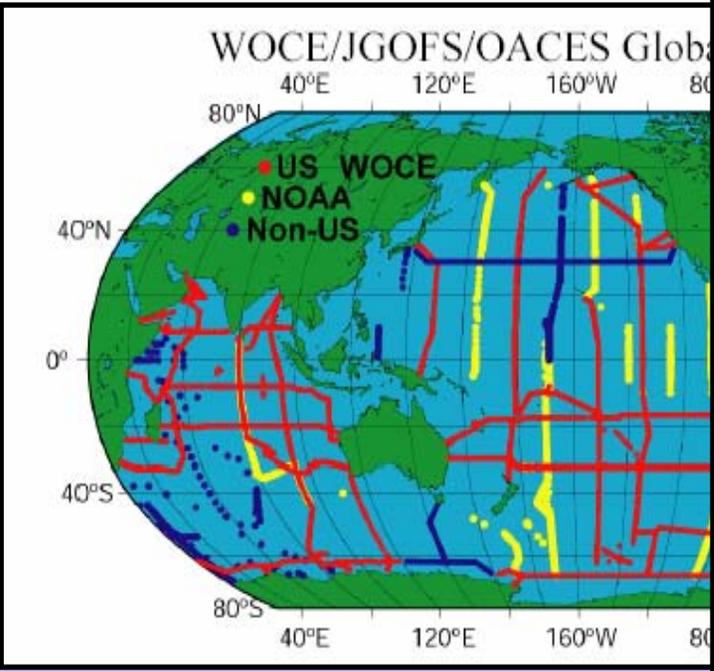
# Decadal Changes in Pacific Ocean Inorganic Carbon

by

Christopher L. Sabine (PMEL), Richard A. Feely (PMEL), Frank Millero (RSMAS), Andrew Dickson (SI O), Chris Langdon (RSMAS), Sabine Mecking (UW), Jim Swift (SI O), Dana Greeley (PMEL)



A first look at the distribution of anthropogenic CO<sub>2</sub> in the ocean was based on the WOCE/JGOFS/OACES global survey of carbon conducted in the 1990s.



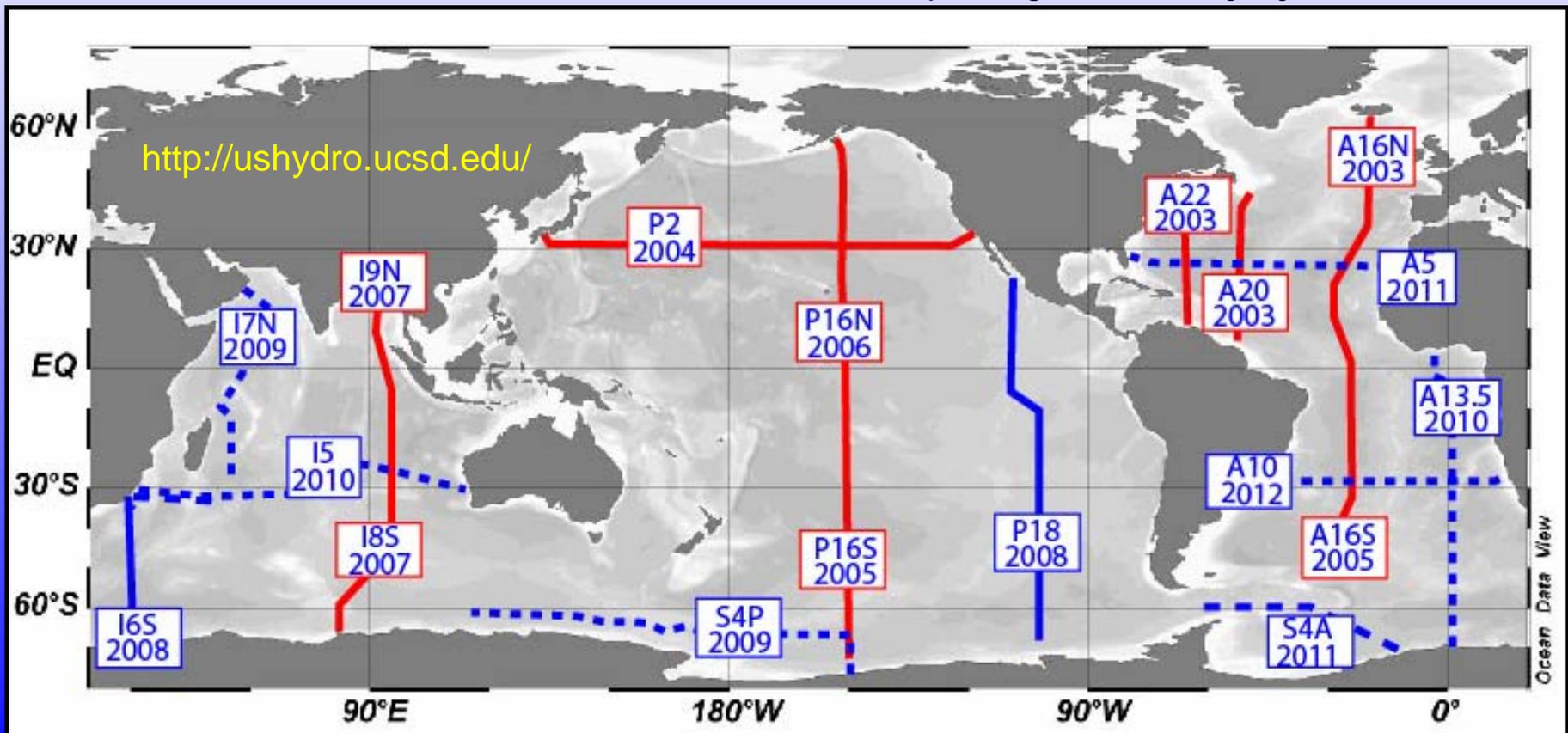
Pacific shows very little penetration in high latitude S.O., deeper penetration in southern subtropics and relatively shallow penetration in tropics

# CLIVAR/CO<sub>2</sub> Repeat Hydrography

**Goal:** To quantify decadal changes in the inventory and transport of heat, fresh water, carbon dioxide (CO<sub>2</sub>), chlorofluorocarbon tracers and related parameters in the oceans.

**Approach:** The sequence and timing of the U.S. CLIVAR/CO<sub>2</sub> Repeat Hydrography cruises have been selected so that there is roughly a decade between them and the WOCE/JGOFS global survey.

**Achievements:** The U.S. CLIVAR/CO<sub>2</sub> Repeat Hydrography Program has completed 9 of 18 lines and is on schedule to complete global survey by 2012.



Comparison of profiles from stations near the intersection of P2 and P16N.

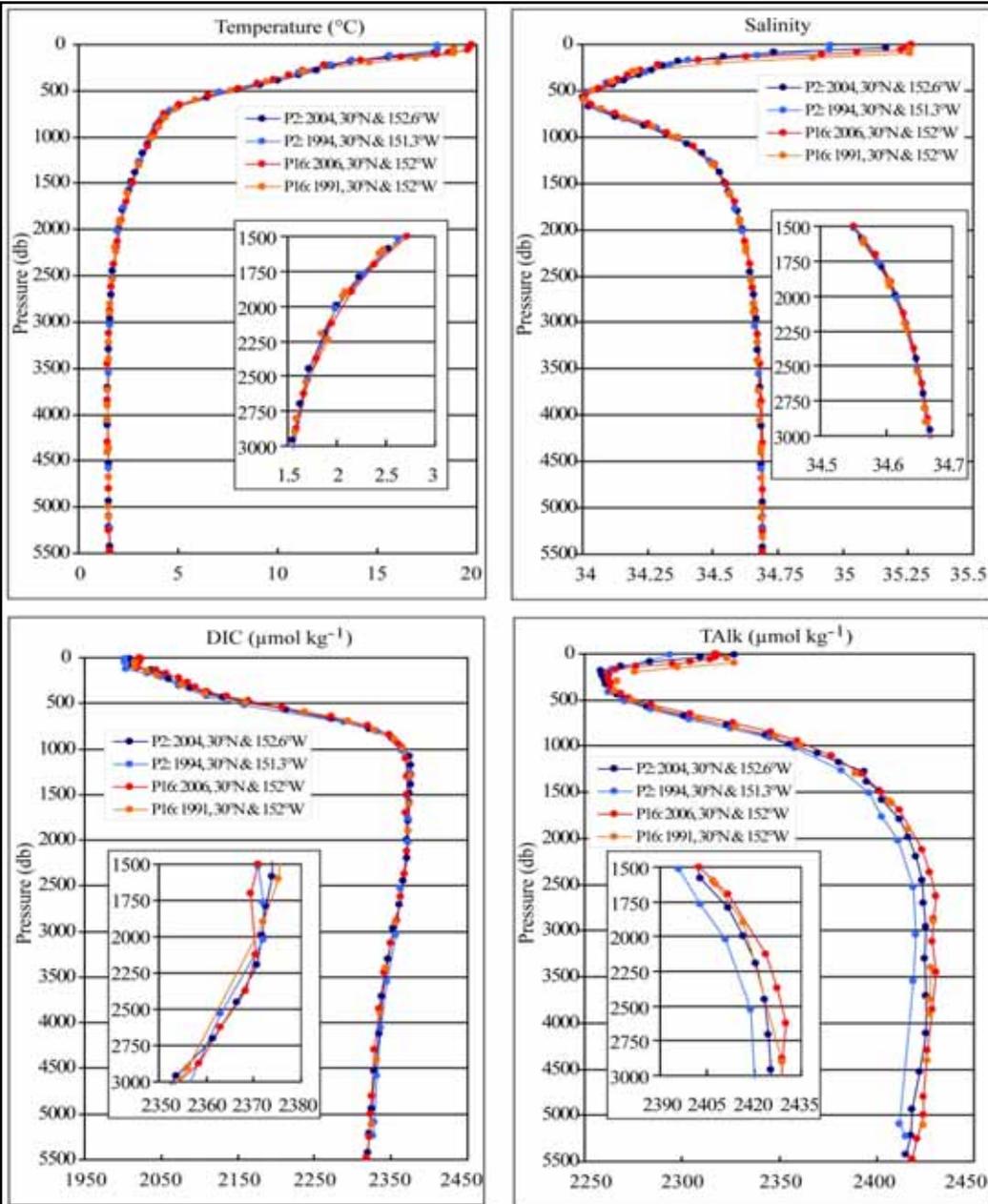
Repeat Hydrography Data Are Very High Quality

P02 along 30° N  
Japan to San Diego, CA  
June-August 2004

P16S along 150° W  
Tahiti to New Zealand  
Jan. - Feb. 2005

P16N along 152° W  
Tahiti to Kodiak, AK  
Feb. - March 2006

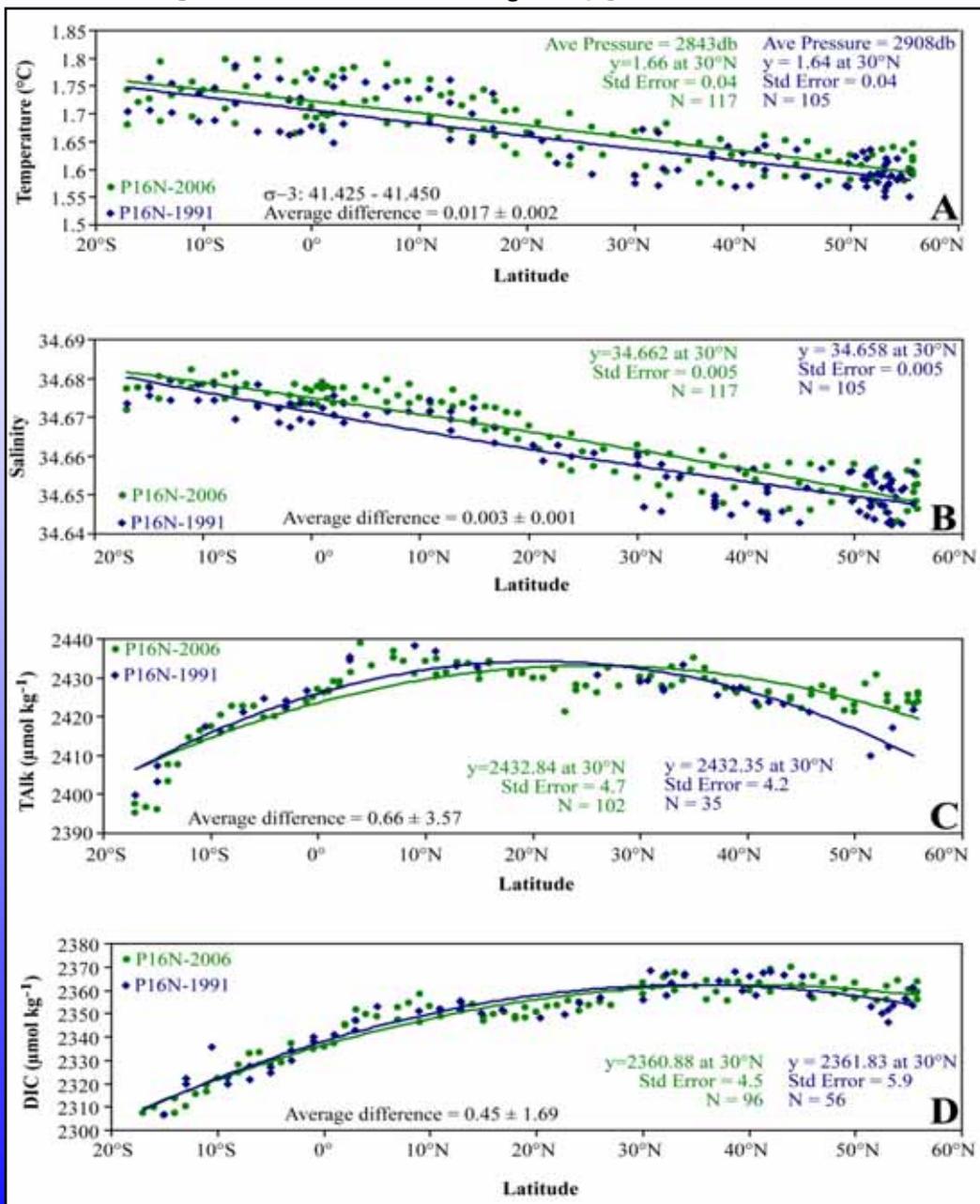
Comparison of crossover and overlap stations indicate the DIC data are good to +/- 1  $\mu\text{mol kg}^{-1}$  and alkalinity data are good to +/- 2  $\mu\text{mol kg}^{-1}$



Comparison of 1991 P16N data with 2006 P16N data along 41.425-41.450  $\sigma_3$  isopycnal surface.

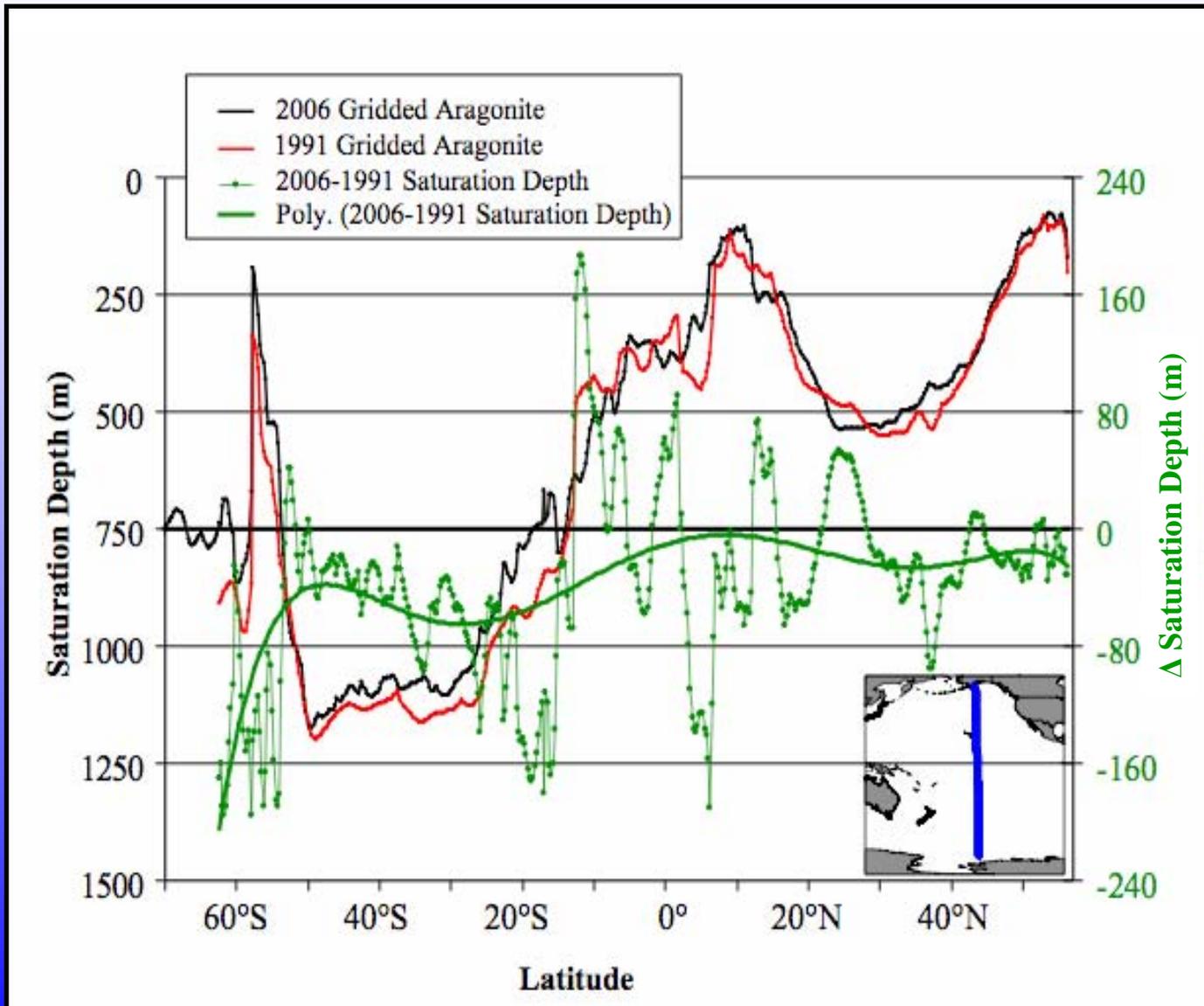
## Repeat Hydrography Data Agree Well With Historical Data

These cruises repeat WOCE lines P02 occupied in 1994 (10 yr diff.)  
 P16S/P16A in 1991/2 (14 yr diff.)  
 P16C/P16N in 1991 (15 yr diff.)



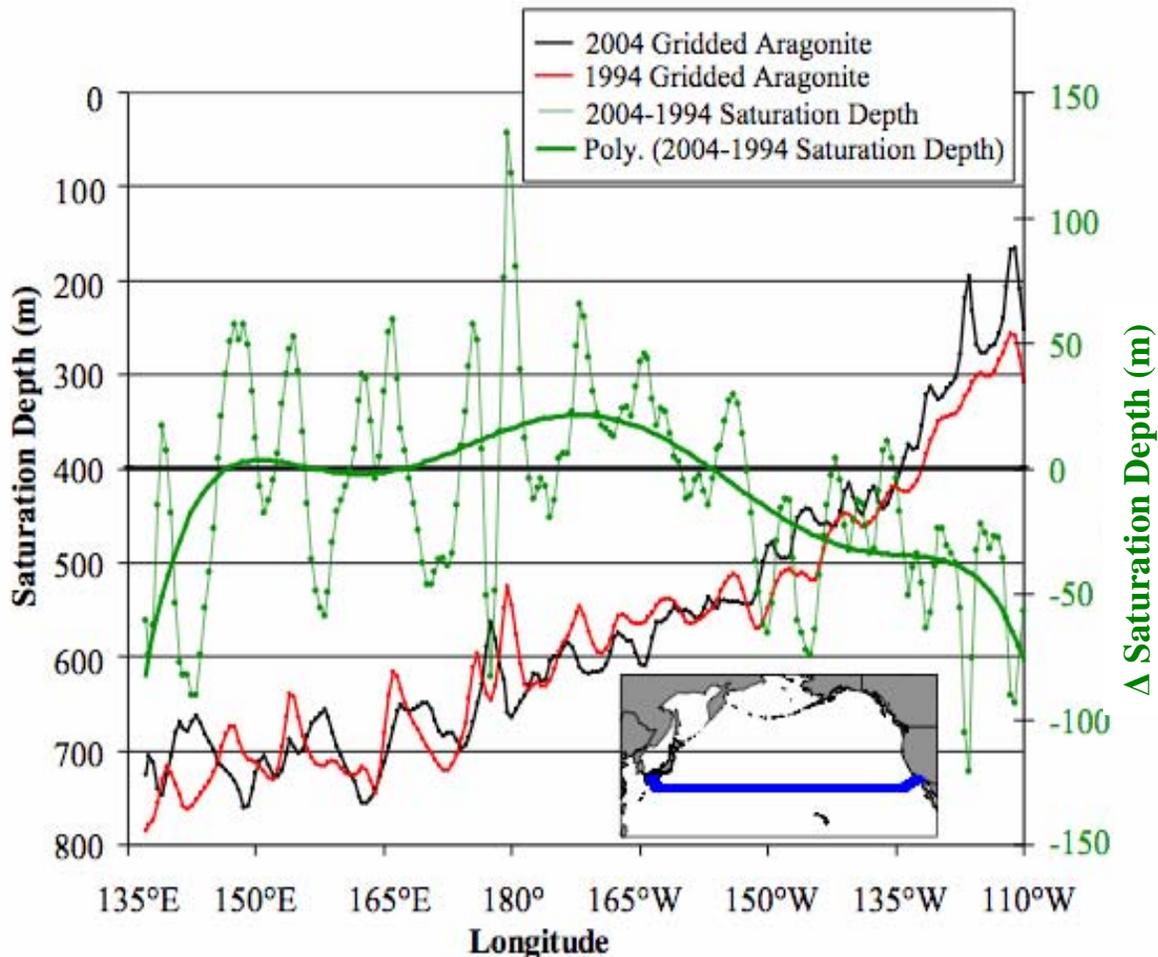
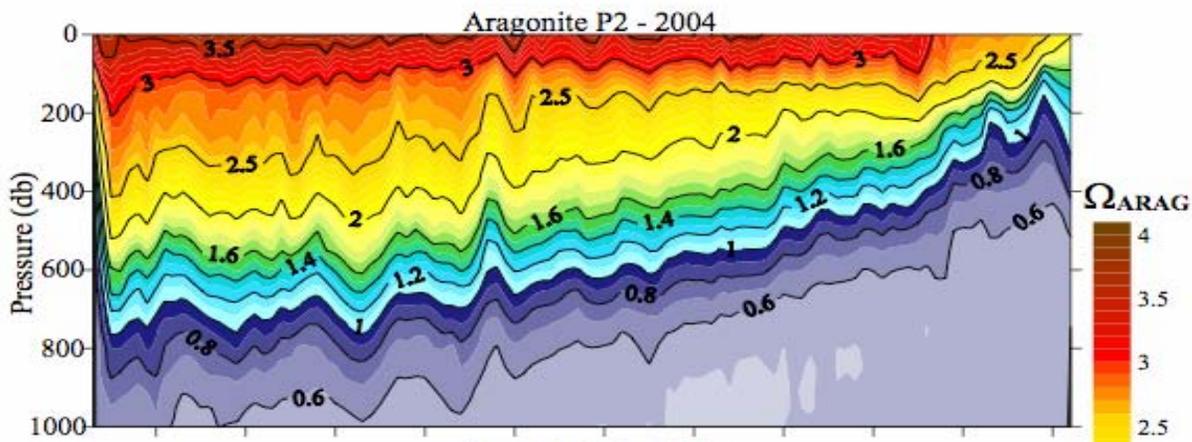
Comparison of deep waters on isopycnal surfaces show no significant offsets between Repeat Hydrography and WOCE cruises.

# Shoaling of aragonite saturation horizon of $\sim 1\text{-}2\text{ m yr}^{-1}$



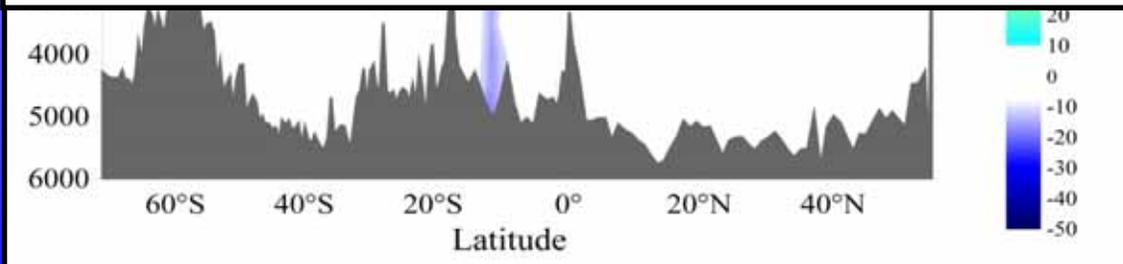
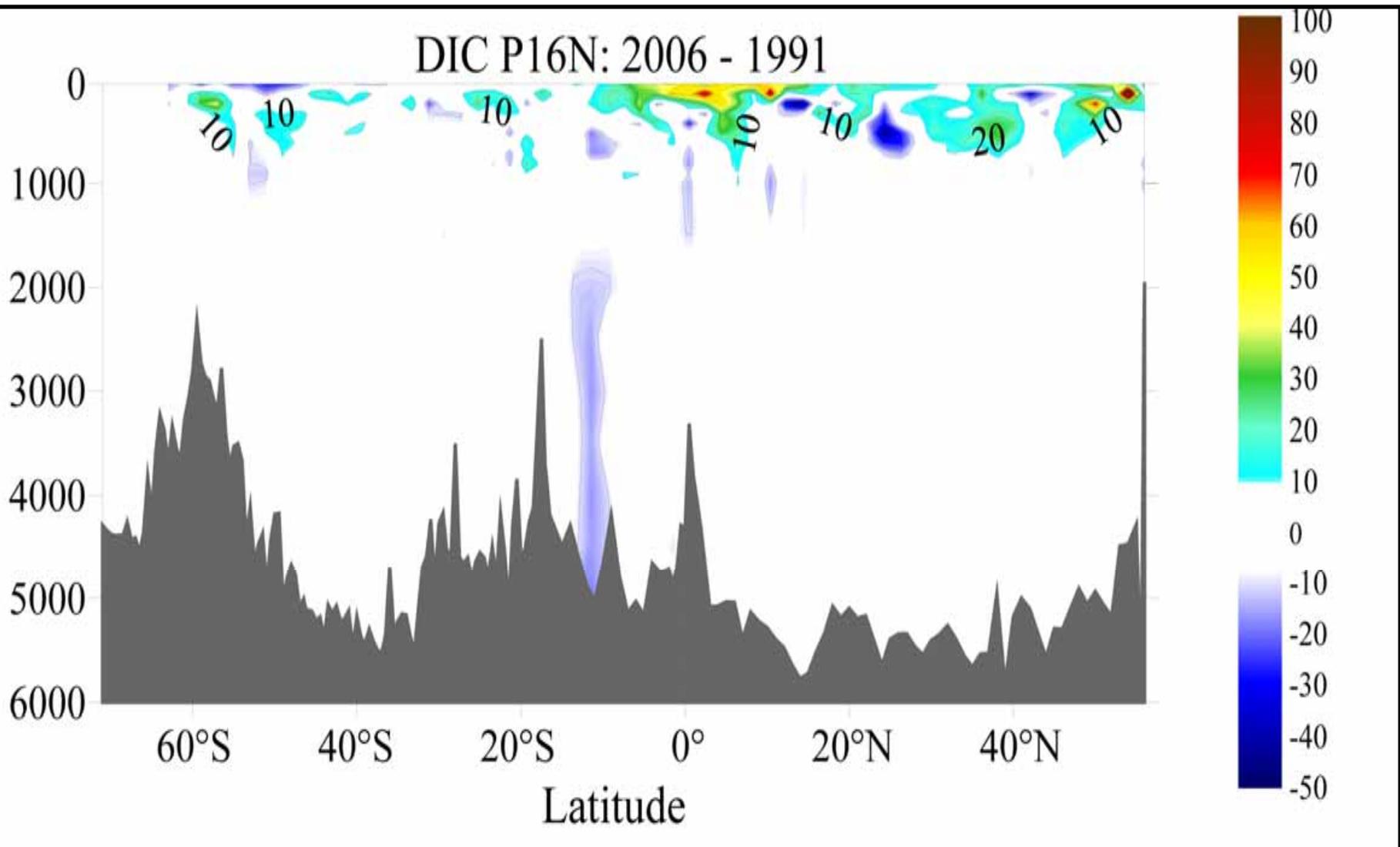
Decreases of aragonite saturation depths in the upper 1000m

Shoaling of aragonite saturation horizon of  $\sim 1\text{-}2\text{ m yr}^{-1}$



Major shoaling of aragonite saturation in the eastern Pacific

How do circulation changes affect decadal carbon chemistry signals?



**But the changes are very patchy and do not show consistent trends.**

## Use a Multiple Linear Regression Approach to Identify Changes

Wallace (1995, OOSDP Report #5) first recognized that empirical relationships between carbon and other hydrographic properties could be used to isolate the CO<sub>2</sub> uptake in the ocean.

Approach:

- 1) Fit carbon data from older cruise with properties that should not be affected by rising atmospheric CO<sub>2</sub>,
- 2) Use empirical fit of older cruise together with hydrographic data from new cruise to predict carbon distributions on the new cruise,
- 3) The difference between the measured carbon values on the new cruise and the predicted values is a measure of the additional carbon taken up from the atmosphere.

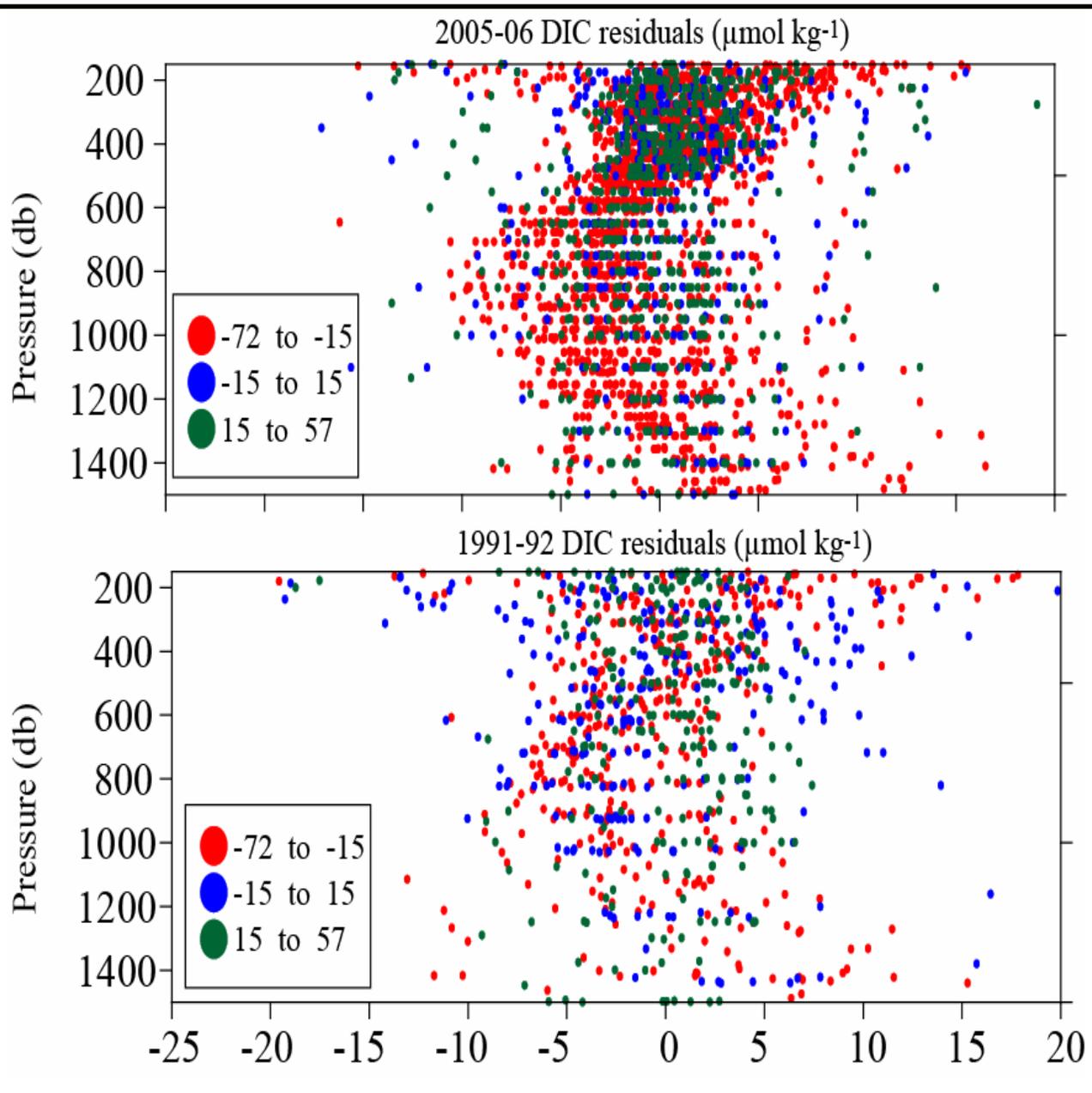
Friis et al. (2005, Deep Sea Res.) refined this approach with the extended MLR where both cruises are fit and take difference in fits.

$$DIC_{(1991)} = a * \sigma_{\theta} + b * \theta + c * S + d * Si + e * P + f$$

$$DIC_{(2006)} = A * \sigma_{\theta} + B * \theta + C * S + D * Si + E * P + F$$

$$\Delta DIC_{(06-91)} = A - a * \sigma_{\theta} + B - b * \theta + C - c * S + D - d * Si + E - e * P + F - f$$

## Use a Multiple Linear Regression Approach to Identify Changes



$$\text{DIC} = f(\sigma_{\theta}, \theta, S, S_i, P)$$

P16 2005/6:

72°S-15°S  $R^2=0.994$ ,  
Std err.=4.7, n=1552

15°S-15°N  $R^2=0.994$ ,  
Std err.=4.4, n=392

15°N-57°N  $R^2=0.999$ ,  
Std err.=4.1, n=575

P16 1991/2:

72°S-15°S  $R^2=0.992$ ,  
Std err.=5.3, n=418

15°S-15°N  $R^2=0.989$ ,  
Std err.=6.6, n=253

15°N-57°N  $R^2=0.999$ ,  
Std err.=3.8, n=250

# Use a Multiple Linear Regression Approach to Identify Changes

$$\text{DIC} = f(\sigma_{\theta}, \theta, S, \text{Si}, P)$$

P2 2004:

135°E-145°E  $R^2=1.000$ ,  
Std err.=2.2, n=199

145°E-140°W  $R^2=1.000$ ,  
Std err.=2.5, n=1389

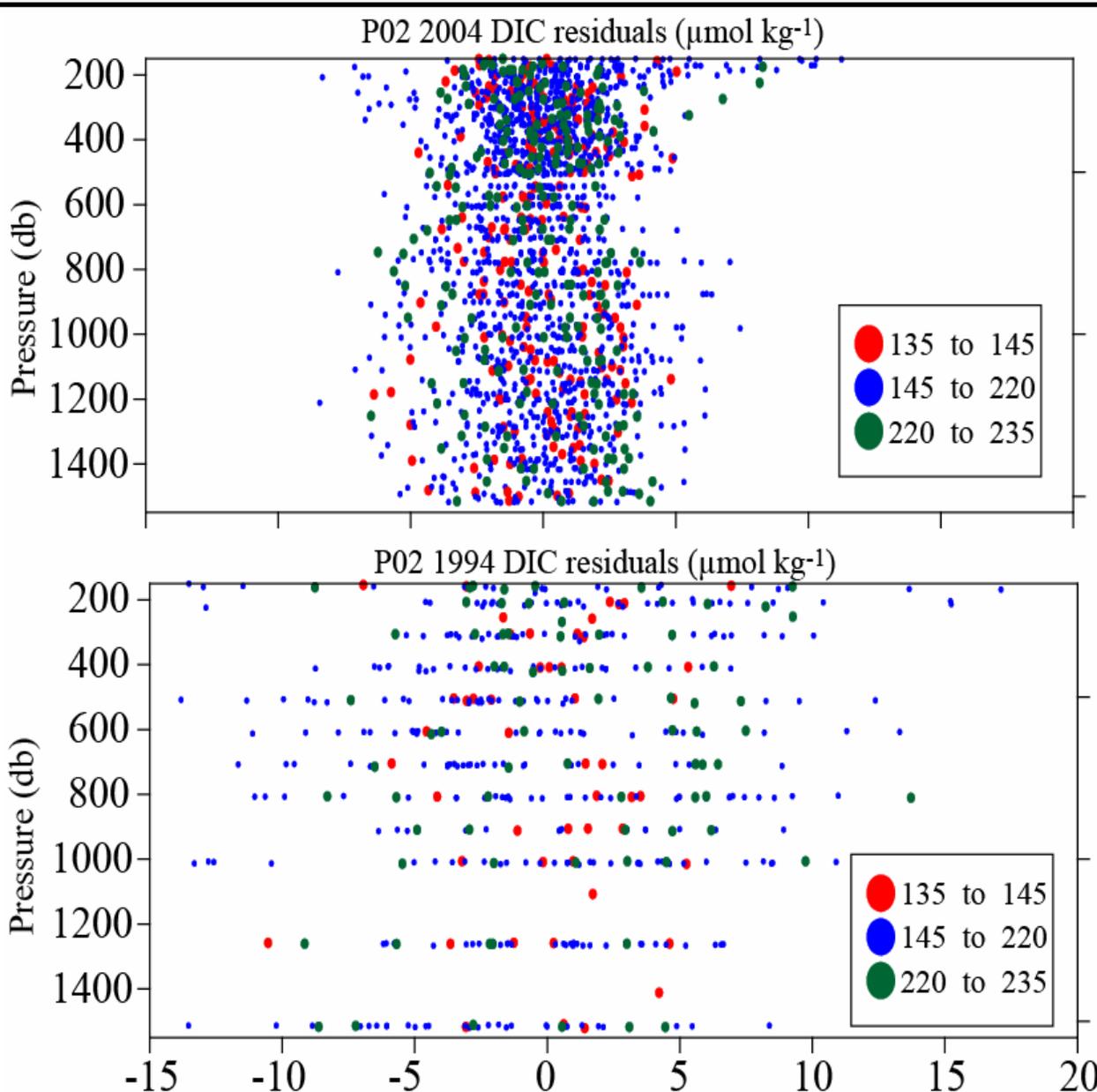
140°W-125°W  $R^2=0.999$ ,  
Std err.=2.5, n=240

P2 1994:

135°E-145°E  $R^2=0.999$ ,  
Std err.=3.6, n=50

145°E-140°W  $R^2=0.998$ ,  
Std err.=5.8, n=353

140°W-125°W  $R^2=0.998$ ,  
Std err.=5.8, n=82



# Total Change in DIC Along P16 (150°W) 2005/6 – 1991/2

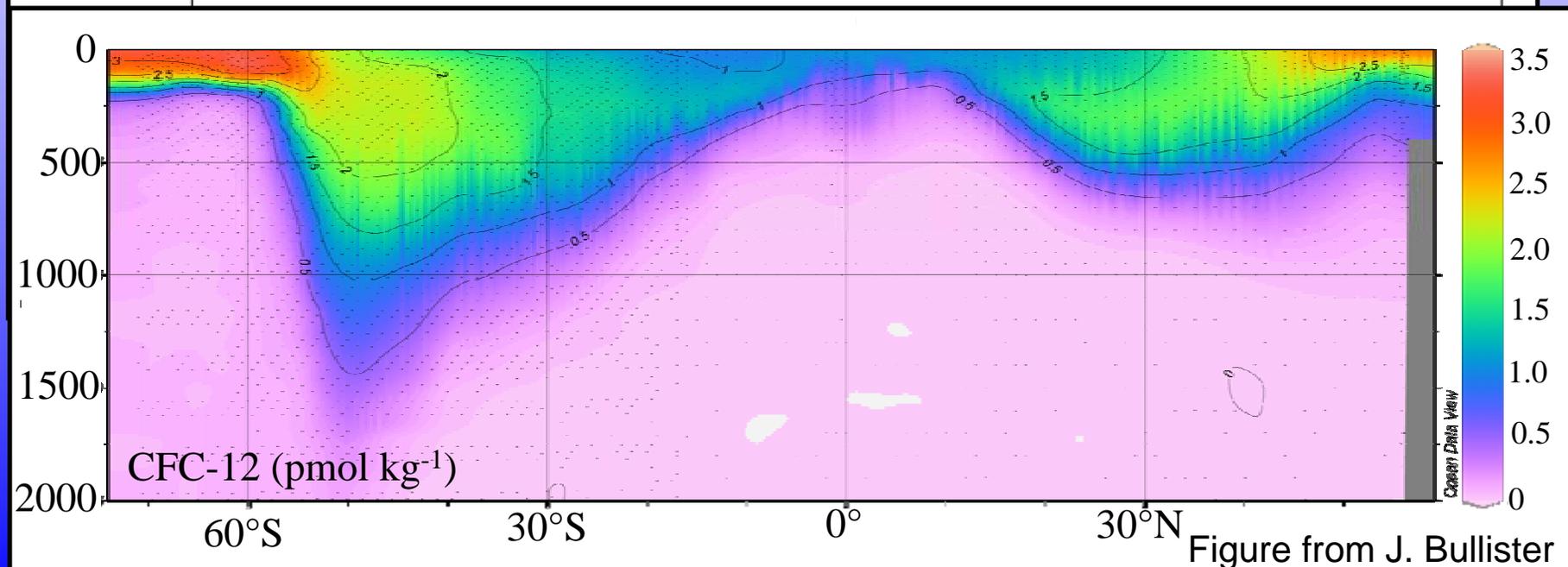
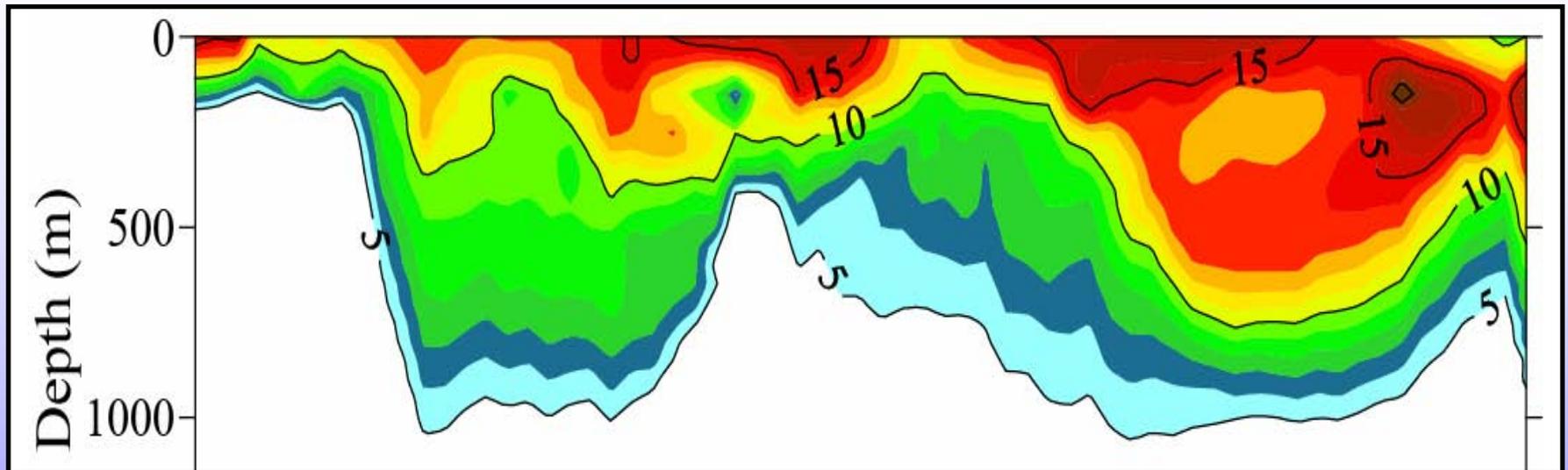
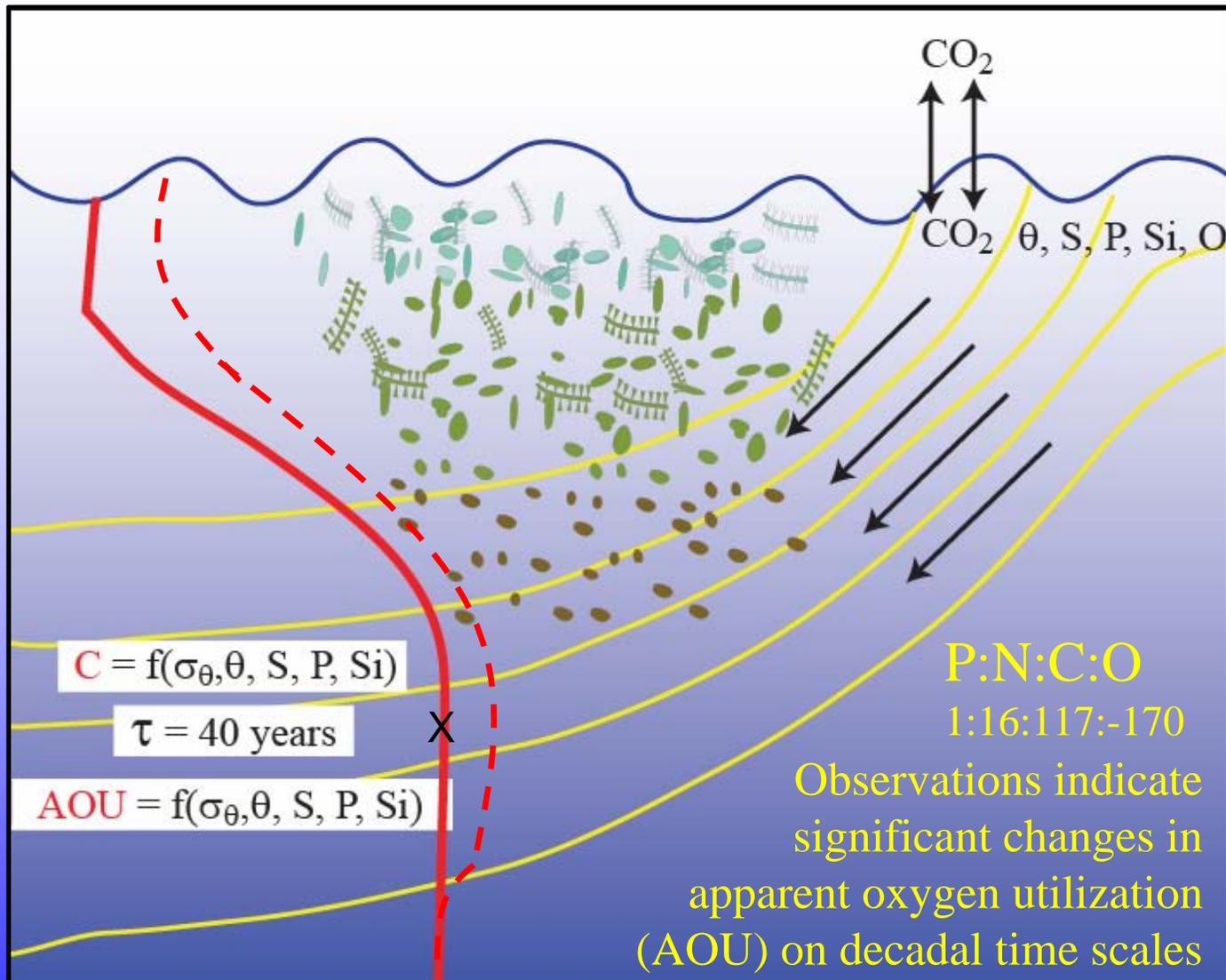


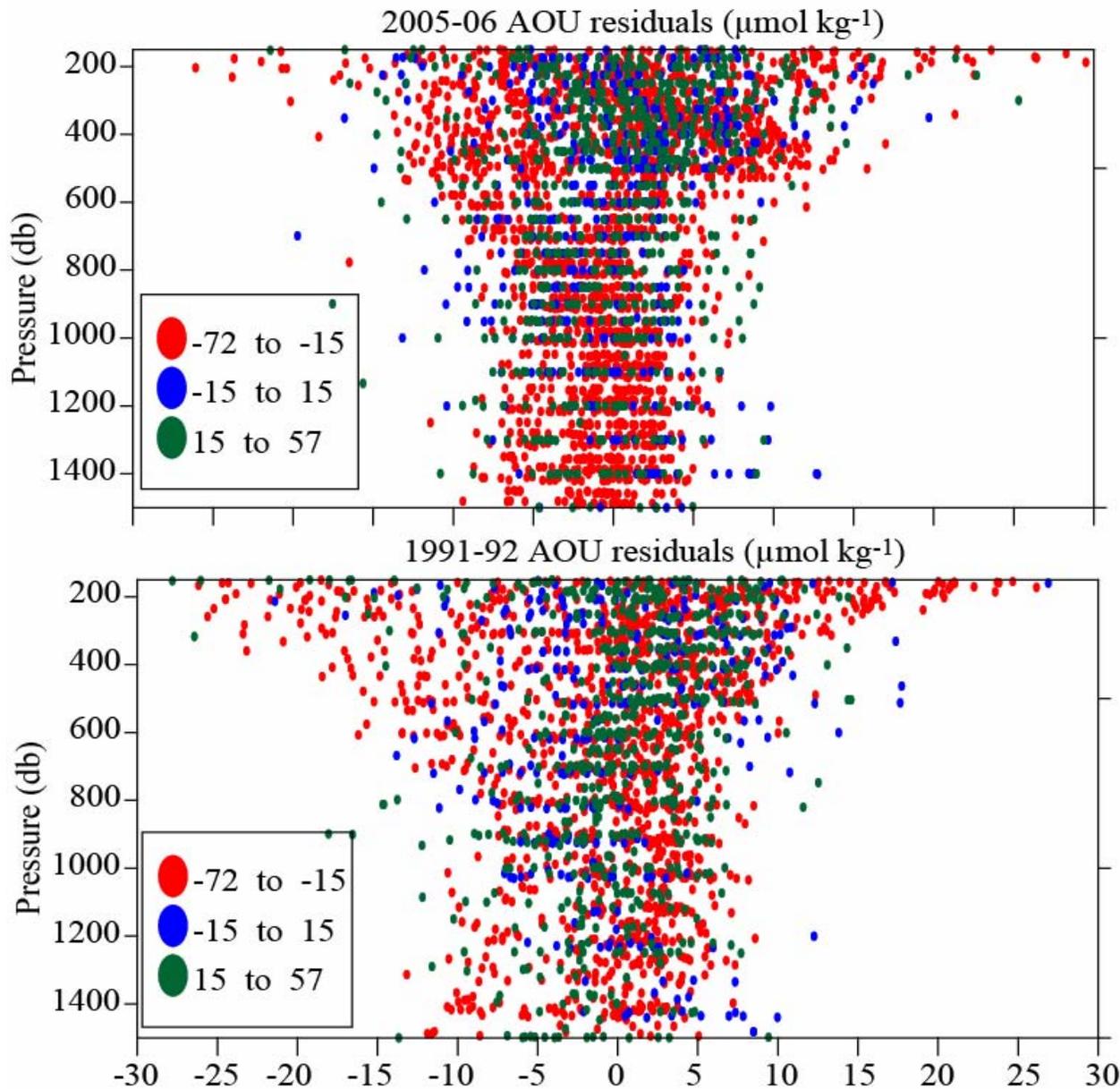
Figure from J. Bullister

# What Does the MLR Tell us About Carbon Changes?



Changes in circulation or changes in export flux can alter the apparent remineralization rate affecting the carbon distributions relative to the other parameters - - use AOU to estimate this change

## Use a Multiple Linear Regression Approach to Identify Changes



$$\text{AOU} = f(\sigma_{\theta}, \theta, S, S_i, P)$$

P16 2005/6:

72°S-15°S  $R^2=0.984$ ,  
Std err.=6.2, n=1829

15°S-15°N  $R^2=0.987$ ,  
Std err.=6.0, n=446

15°N-57°N  $R^2=0.996$ ,  
Std err.=6.1, n=653

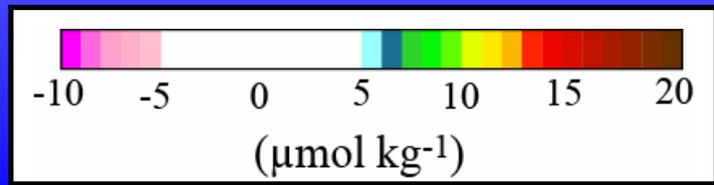
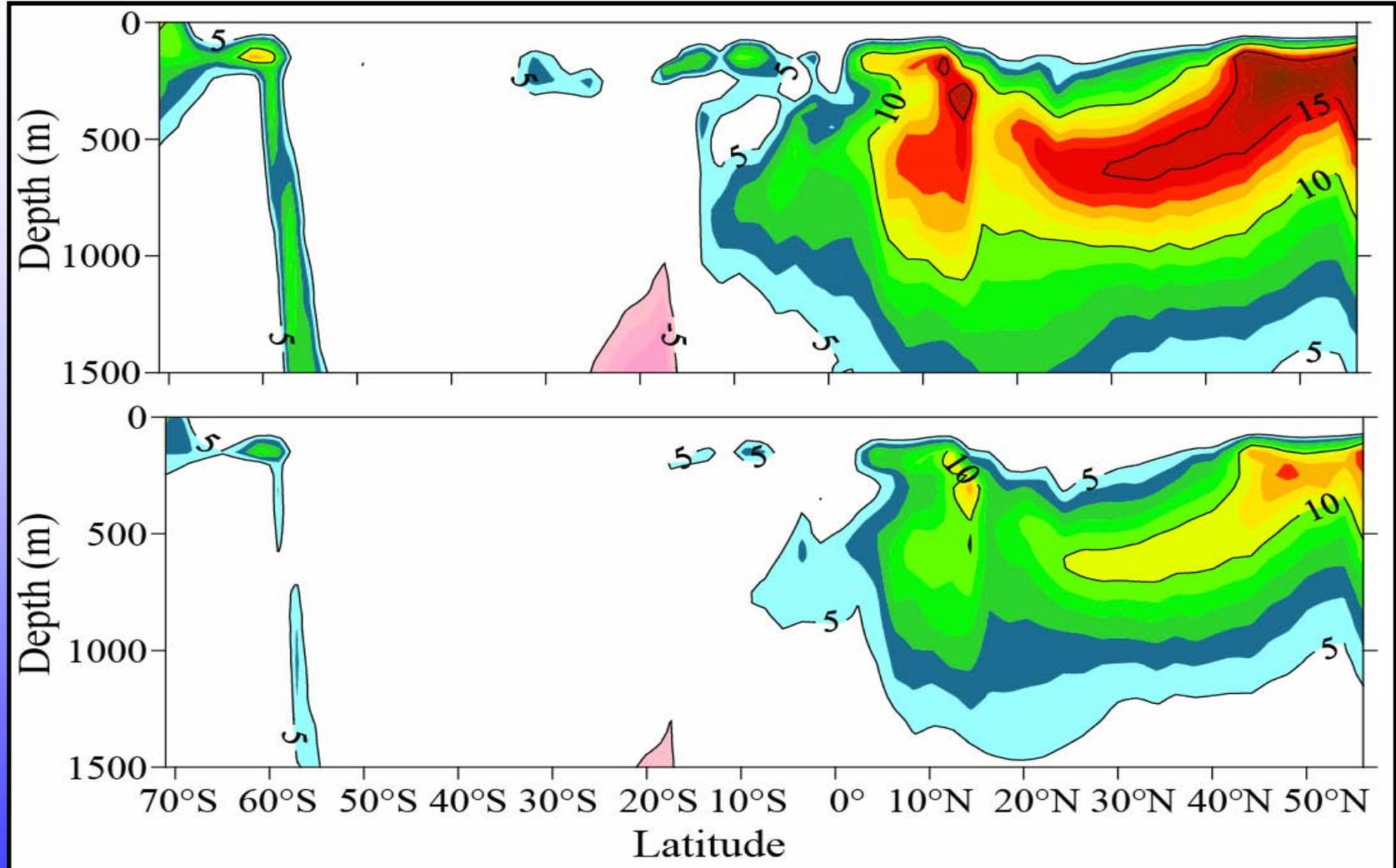
P16 1991/2:

72°S-15°S  $R^2=0.977$ ,  
Std err.=7.4, n=1376

15°S-15°N  $R^2=0.986$ ,  
Std err.=6.4, n=292

15°N-57°N  $R^2=0.995$ ,  
Std err.=6.3, n=655

# Change in AOU (top) And AOU Converted to C Using 117/170 Redfield Ratio (Bottom) Along P16 (150°W) 2005/6 – 1991/2



# Total - AOU DIC Change Compares Well With CFC-12 Distributions Along P16 (150°W) 2005/6 – 1991/2

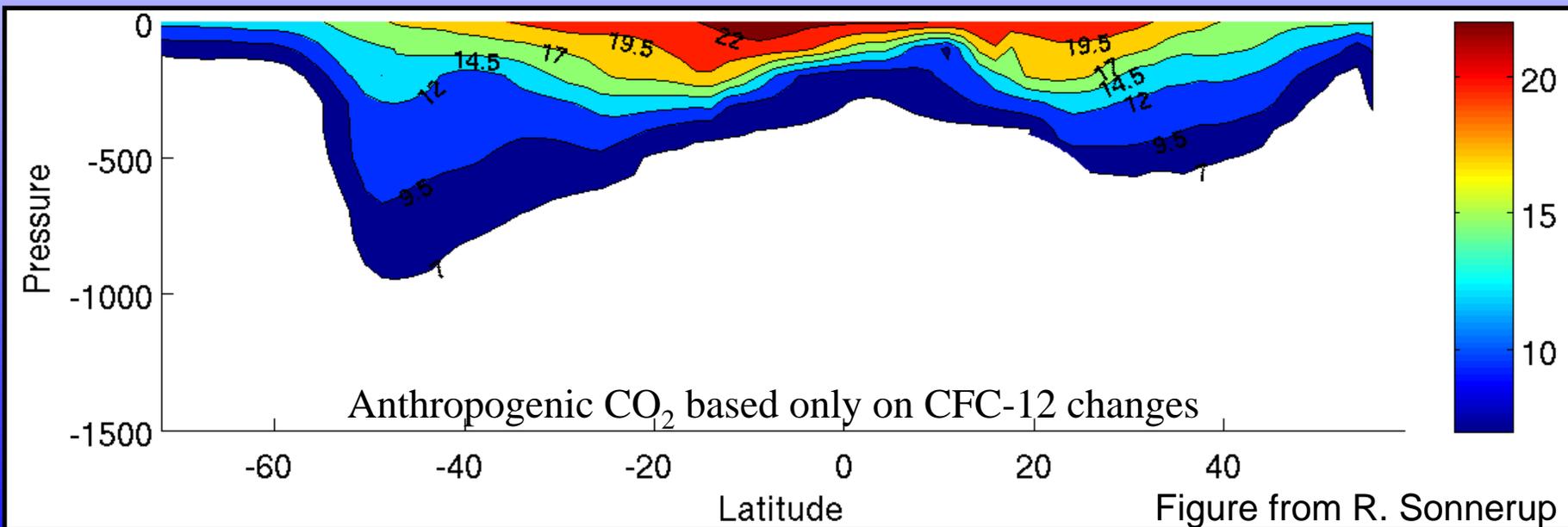
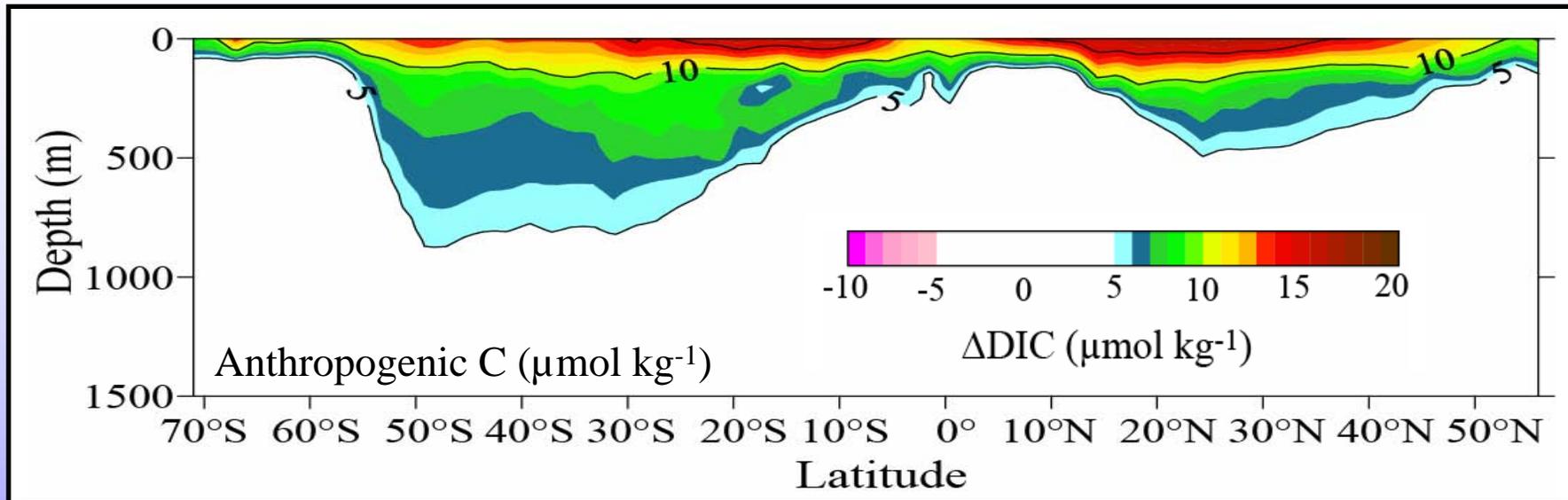
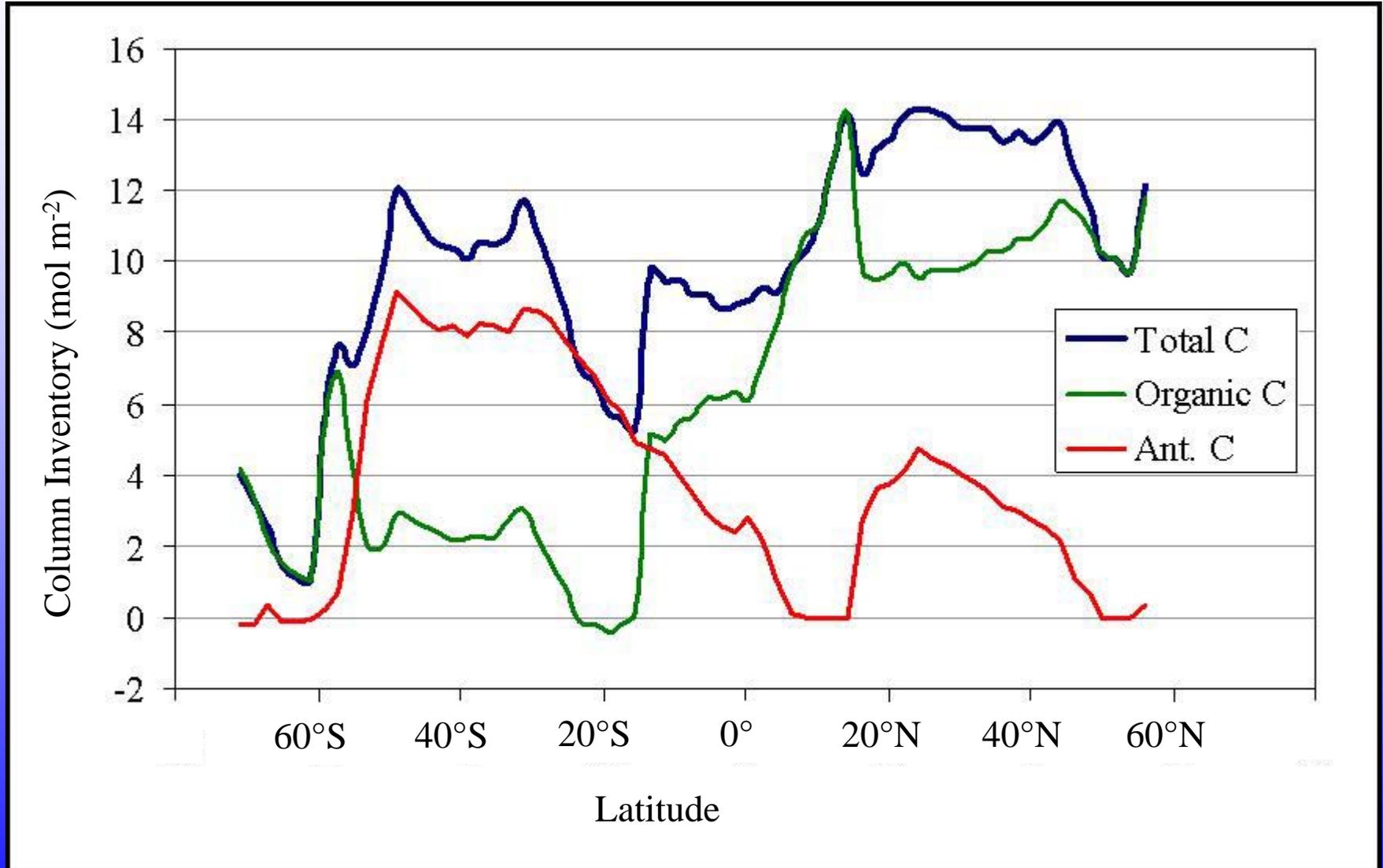


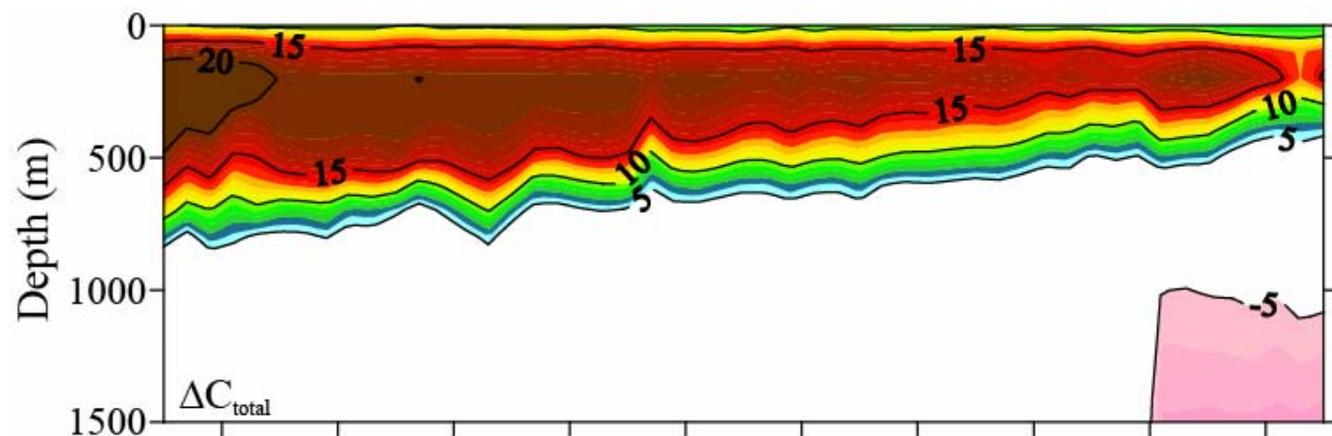
Figure from R. Sonnerup

# Changes in P16 Carbon Inventory Over The Last 14 Years

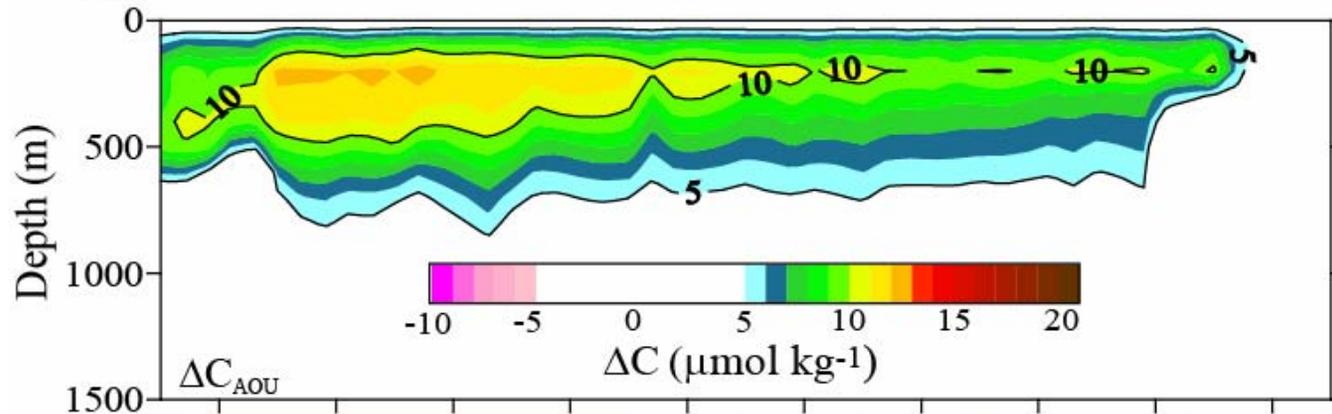


# P02 (30°N) Data 2004 - 1994

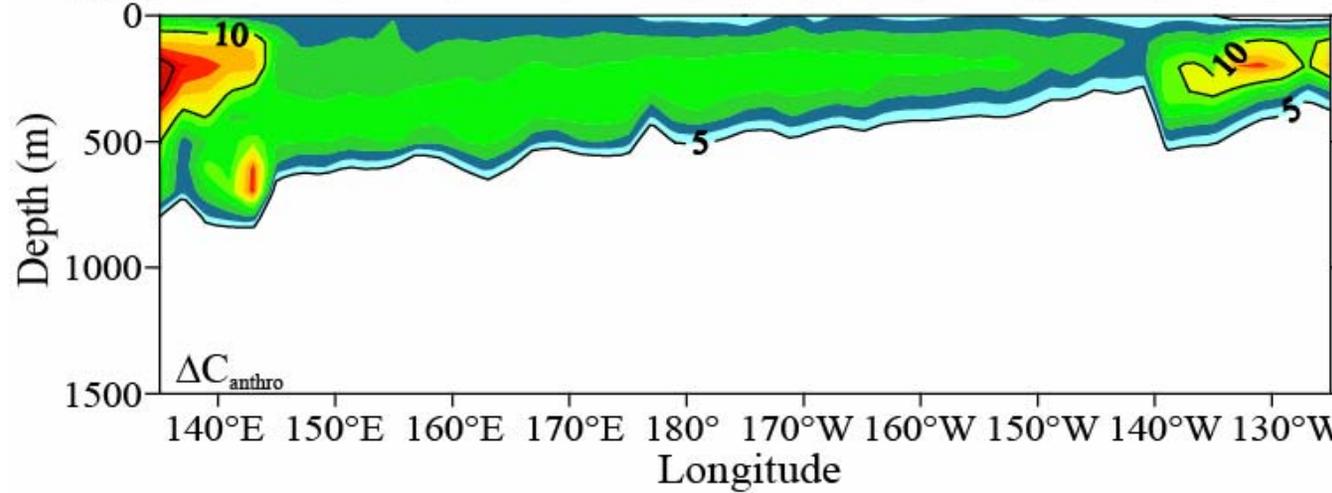
Total Carbon Change



AOU Change  
converted to C  
units



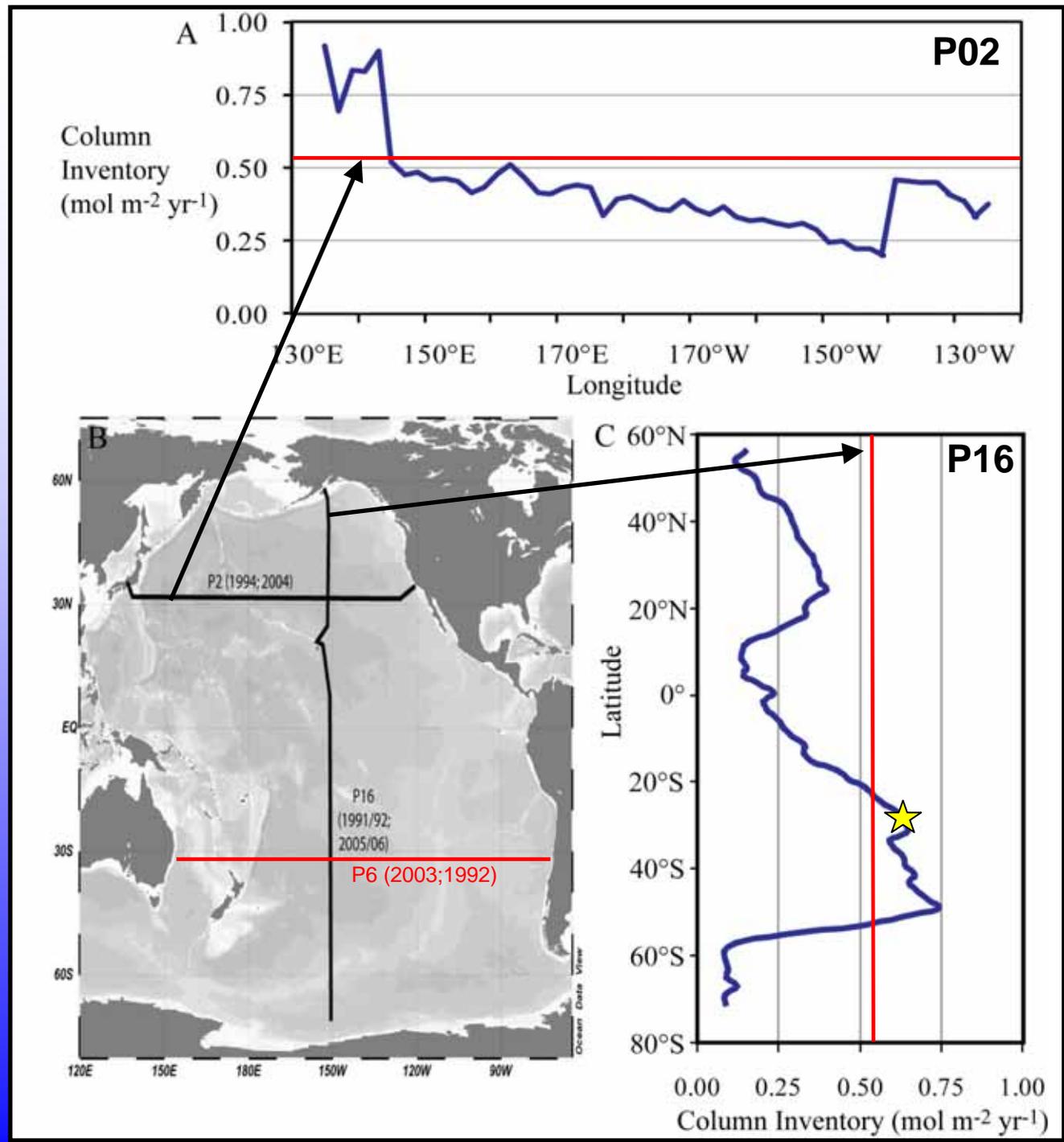
Anthropogenic  
Carbon Change



Average Global Growth Rate of Anthropogenic C is  $0.55 \text{ mol m}^{-2} \text{ yr}^{-1}$  Based on uptake of  $2.2 \text{ Pg C yr}^{-1}$  over a global ocean area of  $335.2 \times 10^9 \text{ km}^2$

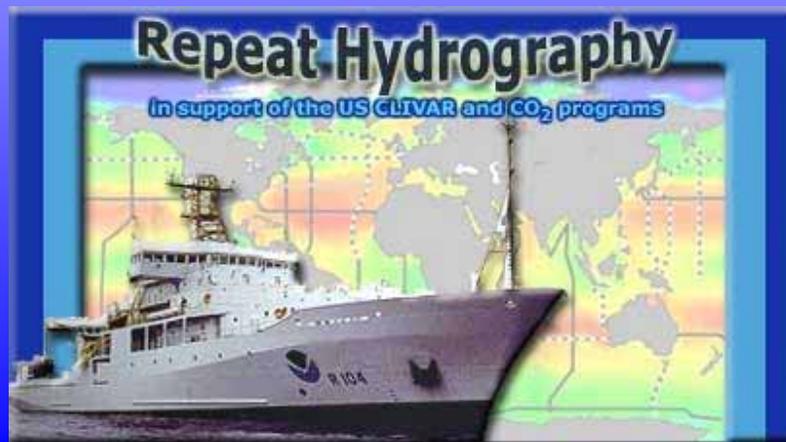


Murata et al. (2007)  
From BEAGLE cruise  
isopycnal analysis  
 $170\text{-}150^\circ\text{W}$



# Conclusions

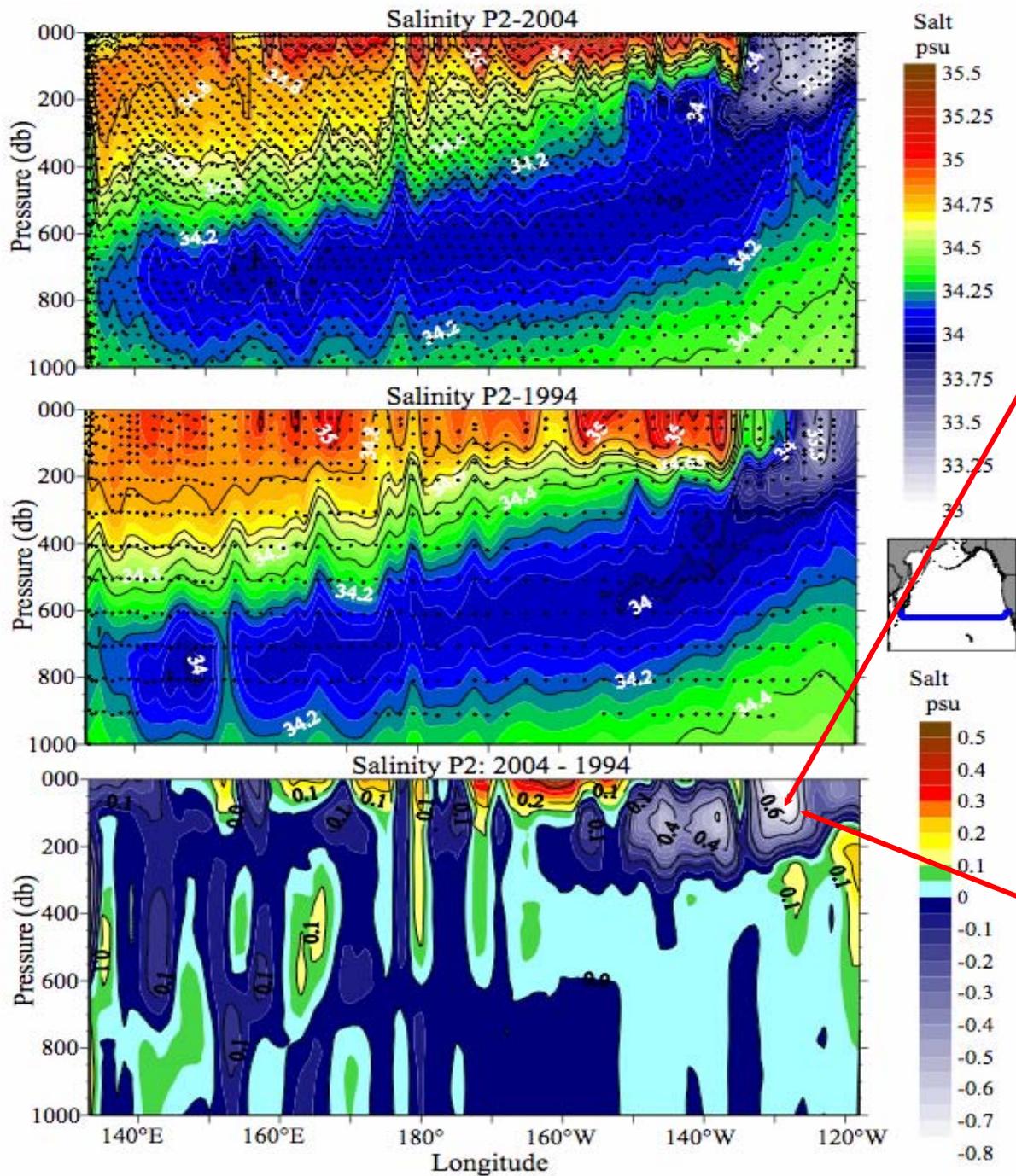
- 1) The Repeat Hydrography program is providing very high quality data
- 2) The observations reveal significant changes in carbon concentrations and aragonite saturation depths on decadal time scales
- 3) Changes in organic remineralization rates can have a significant impact on total carbon changes on decadal time scales
- 4) Both the anthropogenic and organic carbon changes show patterns of variability consistent with expected processes
- 5) Levels of anthropogenic carbon uptake in the Pacific are consistent with anticipated global average uptake



Thank you for your time!



The R/V Thomas G. Thompson arriving in Papeete, Tahiti  
for the beginning of P16N February 2006



# Significant freshening of the California Current in the eastern Pacific

Major decrease in salinity and temperature in the eastern Pacific due to a change in circulation after 1997